Surface Atmosphere Radiation Budget (SARB) working group update

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Outline of this presentation

- Edition 4 data products
	- CRS
	- EBAF (Edition 4.2)
	- MATCH (NOAA20 aerosol optical thickness consistent with Aqua)
- Edition 5 algorithm developments
	- Aerosol transport model
	- CERES radiative transfer model
	- Polar surface albedo

Edition 4 products

- Edition 4 CRS
	- Instantaneous surface and in atmosphere irradiances (Level 2)
	- Gridded instantaneous irradiances (SYN1deg, Level 3)
	- January 2018 through December 2022 (Terra or Aqua) were released
- Edition 4.1 SYN
	- Gridded (1 deg × 1deg) hourly, daily, and monthly mean surface and in atmosphere irradiances.
	- Produced through June 2024.
	- Edition 4B (MET-10 bug fix, 2-channel nighttime optical depth, and no twilight algorithm with TISA new interpolation) available soon.
- Edition 4.2 EBAF
	- Gridded (1 deg × 1deg) monthly mean surface irradiances
	- Produced through May 2024
	- Reprocess clouds over the NOAA20 period (April 2022 onward) with MERRA-2 .
	- Climatological adjustments to NOAA20 will be applied with Terra+Aqua using a common period from May 2018 through March 2022.
	- The revised product (Edition 4.2.1) will be released early 2025
- Edition 4 MATCH
	- Consistency of aerosol optical thickness derived from MODIS and VIIRS
	- Working with the Deep Blue team to mitigate the AOD differences
- CCCM D2 version
	- Produced with CALIPSO V4-51and CloudSat R05 data products.
	- Will be released soon.

Edition 4.2 EBAF climatological adjustment in Wm-2

Kato et al. 2024

Terra, Terra+Aqua, and NOAA20 Global monthly anomaly time series Blue: with climatological adjustments Red: without climatological adjustment

Anomaly (Wm⁻² -2 $0²$ Year

b) Downward longwave irradiance

a) Downward shortwave irradiance

Edition 4 aerosol (MATCH)

- Because the Deep Blue algorithm uses a newer algorithm for NOAA20 VIIRS than Terra and Aqua, there are significant discontinuity in aerosol optical thickness.
	- Jaehwa Lee provided coefficients to correct NOAA20 AOT.
- There are some discontinuities of Dark Target aerosol optical thickness over ocean and land.
	- We developed coefficients to correct NOAA20 Dark target optical thickness over land.

Edition 5 aerosol transport model

BC: Black Carbon, POM: Primary Organic Matter, SOA: Secondary Organic Matter, SU: Sulfate, DU: Dust, SS: Sea Salt

CAM6-CERES radiative transfer model interphase

3D and temporal space of mixing ratio of all aerosol types from CAM6 or GEIS-IT

CAM6 are separated by size Accumulation Aitken Coarse Primary Carbon

3D space of Tau, omega, and g as a function of time and wavelength

Tau (surface layer @650 nm) Omega (surface layer @650 nm) g (surface layer @650 nm)

CERES Cloud Radiative Swath (CRS) Update

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> > **Collaboration with:**

TISA Group: Joshua C. Wilkins, David Doelling and Pamela Mlynczak (TISA gridding for CRS1deg product)

Data Management: Walter Miller, Victor Sothcott, Joshua C. Wilkins, and Kathleen Dejwakh

ADM Group: Lusheng Liang, Wenying Su (TOA fluxes)

Cloud group: Bill Smith Jr, Sunny Sun-Mack, and Ben Scarino (Cloud and skin temperature retrievals)

FLASHFLUX group: Paul Stackhouse (Parameterized surface fluxes in FLASHFLUX)

Level 2 CRS Flux Algorithm

CATM : CERES Atmospheric Transport Model produced by CERES SARB group

Changes from the Ed4 to Ed5 CRS Algorithm

Changes from the Ed4 to Ed5 CRS Algorithm

Validation of the CRS Flux Algorithm over Polar Regions

- More frequent observations by the sun-synchronous (Aqua or Terra) satellite orbits over the polar regions, compared to other lower latitude regions (but only daytime during summer and nighttime during wintertime).
- Ability to validate the computed surface radiation budget using ground site measurements with a diurnal cycle
- The surface type is usually snow and sea ice (SIC), meaning that the assumption of high surface albedo is important in the radiative transfer.
- Uncertainties in the skin temperature of the analysis dataset are large over the polar regions, requiring further examination of the input temperature and humidity properties to the radiative transfer model.

"Permanent snow" "Permanent snow" "Partly or completely Sea Ice (SIC)"

Constraining Surface Spectral Albedo Using Observed Surface BB albedo

- First guess of the spectral surface albedo is based on the following sources.
	- \checkmark MODIS land surface BRDF product over land (MCD43C1)
	- \checkmark Jin's ocean, snow, and sea ice (SIC) spectral albedo model (Jin et al., 2004, 2008)
- These spectral albedos are constrained by the observational-based broadband (BB) surface albedo, by deriving a scaling factor. The observed BB surface albedo is from TOA (CERES or MODIS) observations for clear sky cases, called **S**urface **A**lbedo **H**istory (SAH) Map.
- Until Ed4, the scaling factor was derived regardless of the cloudy conditions. This caused problems for cloudy skies since the observed BB albedo is only from clear skies. The cloudy-sky snow/SIC BB albedo is larger than the clear-sky albedo for snow or sea ice (SIC) surface types. In the Ed5 algorithm, we derive the scaling factor based on the clear sky assumption. Then the scaling factor is used for total, clear, no-aerosol, and pristine skies simulations.

Better Angular Correction of the Observed Snow Surface Albedo Related to the Solar Zenith Angle (SZA) Changes in the Ed5 Algorithm

The surface albedo decreases with increasing cos(SZA). As the value of *d* is larger, a larger decrease of the surface albedo over cos(SZA).

- The monthly gridded Surface Albedo History (SAH) map provides the observed surface BB albedo at a certain solar zenith angle (SZA).
- The angular correction is needed to get the surface albedo at the desired SZA. In the CRS algorithm, the angular correction model is based on Dickinson's diurnal variation model of the albedo (1983): $\alpha(\mu_2) = \alpha(\mu_1)$ $(1 + 2d\mu_1)$

 A smaller d value means a smaller variation of the albedo over the SZA.

 $(1 + 2d\mu_2)$

• For the snow surface type, d=0.1 was used for the Ed4 CRS algorithm. In the Ed5 algorithm, d=0.05 is used. This is more consistent with Jin's snow spectral albedo models and are also with observations (next slides).

Better Angular Correction for the Snow Surface Albedo Related to the Solar Zenith Angle (SZA) Changes

Biases of the CRS SW TOA fluxes with d=0.1 over snow Biases of the CRS SW TOA fluxes with d=0.05 over snow

Two Antarctic Ground Sites

- CERES-derived observed TOA fluxes are compared with CRS computed TOA fluxes, as a function of cos(SZA) over the two Antarctic ground sites.
- The new angular correction of the surface albedo (red dots) gives better agreements of the TOA fluxes with the CERES observations, and the TOA biases are less dependent on the SZA, compared to the old angular correction method (black dots).

Improving Biases in the Snow Albedo for Cloudy Skies

Dec 2019 SDM Albedo vs time $1.0 \sqsubseteq$ **Observation** • Original CRS • New CRS **Surface Albedo** 0.9 0.7 0.6 200 600 400 Hour **Jan 2020 SDM Albedo vs time** 1.0 Surface Albedo 0.7 0.6 Ω 200 400 600 Hour

Comparison of Surface Albedo over the Siple Dome Ground Site

- One of issues in the Ed4 algorithm was the underestimation of the snow surface albedo for cloudy skies.
- The underestimation was related to the scaling factor derived regardless of the cloud conditions.
- The new albedo scaling factor is derived in the Ed5 algorithm based on the clear-sky assumption, reducing the underestimation issues.

Ground Observation

Original Ed4 CRS snow albedo algorithm (scaling factor was derived regardless of cloudy conditions) Modified Ed4 CRS snow albedo algorithm (scaling factor is based on the clear-sky assumption)

Improving Biases in the Sea Ice (SIC) Surface Albedo for Cloudy Skies

• Still the broadband albedo is improved by using the new observational constraining method over the Arctic.

The Ed5 Fu-Liou Model

- \square Main radiative solver (4-stream for SW and 2-stream for LW) remains the same as in the Ed4 Fu-Liou model
- \Box More flexibility in changing the Fu-Liou band structure
- □ Updating the line-by-line (LBL) dataset (especially water vapor continuum) (Hogan et al. 2020, 2022) and including more gas species in generating CKD table \Box CO₂ and CH₄ are variables for both SW and LW calculations.
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- \Box Aerosol interface to use CAM6 output is under development.

SW Band Structure

Ed4

 \Box Changes in the SW band structures (18 to 29 bands):

- \checkmark To reduce errors due to the correlated-k distribution (CKD) assumptions
- \checkmark To have band boundaries at 400 and 700 nm for the Libera study (visible & split SW bands)
- \checkmark Better optimization method in assigning optimal k terms (still progress)

Ed5

In the Ed5 model, nine gas species $(O_3, O_2, N_2O, N_2,$ CO_2 , H₂O, CH₄,

CFC11, and CFC12) are considered for all Fu-Liou bands.

LW Band Structure (Ed4 & Ed5)

- Ed5 LW band structure remains the same as in the Ed5. However, more gas species are included.
- Red fonts are gas species considered in the Ed4.
- Ed5 considers 9 gas species for all bands.

Changes in Downward Flux by Including More Gas Species $(H₂O, O₃ \rightarrow 9$ Species) for 50 Evaluation Cases

Impact of Inclusion of More Gas Species on the SW Fluxes

- \circ In Ed4, O₂ absorption was not included in most SW Fu-Liou bands. However, $O₂$ absorption can be comparable to O_3 and H_2O absorption in some bands.
- o For UV bands, the impact of ignoring $O₂$ is small since the solar radiation is mostly absorbed in the stratosphere (not shown).
- The impact of ignoring $O₂$ is noticeable near surface in the visible bands (0.59 – 0.79 μm). The difference in surface downward fluxes can change by 1 W m^{-2} .

961 hPa

- Gas absorption for the 1.9-2.5 μm spectral region in the Ed4 model was underestimated by 1 W m-2, consisting of 50% of SW broadband error.
- The underestimation happened since the "correlated k assumption" did not work well for the broad spectral region, where the relative impacts of $CO₂$, $CH₄$, and H_2O absorptions vary by pressure. This problem is more serious for the premix approach used in Ed5.
- Remedies
	- \triangleright Split the NIR spectral region into more bands to hold the correlated k assumptions
	- \triangleright Increase the total k terms for the NIR spectral region

Impact of the Splitting Near-Infrared (NIR) (1.9–2.5 μm) Bands on Flux Calculations

Band 19 (1.90-2.28 μm) nk=5 Band 20 (2.28-2.50 μm) nk=3

 $#$ of bands: 2, Total nk = 8

Case1: Two bands with 8 k terms Case2: Seven bands with 14 k terms

Band 19 (1.90 -1.94 μm) nk=2 Band 20 (1.94 -1.99 μm) nk=2 Band 21 (1.99 - 2.03 μm) nk=2 Band 22 (2.03 - 2.08 μm) nk=2 Band 23 (2.08 - 2.13 μm) nk=2 Band 24 (2.13 - 2.28 μm) nk=1 Band 25 (2.28 - 2.50 μm) nk=3

of bands: 7, Total $nk = 14$

Biases in SW downward fluxes from the CKD method to the LBL results

(Tested with 50 evaluation cases of the CKDMIP study (Hogan et al. 2020))

Ed4 Fu-Liou CKD

– Mean Bias ◼ **RMSE**

• Biases in the downward fluxes for the separated bands are larger in the Ed4 results compared to the Ed5 results, but these are largely cancelled out in the broadband flux biases.

SW Broadband Biases by the CKD Methods Compared to the LBL Results

LW Broadband Biases by the CKD Methods Compared to the LBL Results

Broadband Flux Biases to LBL (50 Evaluation Cases by Hogan et al.)

Ed4 Ed5

Summary

- VIIRS Deep Blue correction was developed by the deep blue team (Jaehwa Lee) and Dark target correction was also developed to be used in Ed4 process.
- Edition 4.2 EBAF-Surface from April 2022 will be reprocessed and released in early 2025 (Edition 4.2.1).
- Edition 5 aerosol transport model and interphase to radiation transfer model continues.
- The snow or sea ice (SIC) albedo was improved by using the scaling factor based on the clearsky assumption. By using the new approach, underestimation of the surface albedo issue was largely removed, showing a better agreement with ground observations.
- The Ed5 Fu-Liou model is under development. This include:
	- \checkmark More flexible interface in case we need changes in the SW or LW band structures, cloud scattering parameter databases, and aerosol scattering databases.
	- \checkmark Updated correlated-k distribution table based on more recent version of line-by-line gas database and more inclusion of gas species
- Preliminary results for the clear-sky conditions show that
	- \checkmark SW and LW BB fluxes for Ed4 and Ed5 are not very different, < 2 W m⁻² differences.
	- \checkmark When the fluxes are compared for narrow bands, the Ed4 and Ed5 are quite different. When comparing with the line-by-line results, Ed5 results show better performance than the Ed4 model. 29

Publications

- Ham, S.-H., N. G. Loeb, S. Kato, T. J. Thorsen, A. Voigt, W. L. Smith Jr., and D. Winker, 2023: Zonal cloud trends observed by passive MODIS and active CALIPO and CPR sensors, Submitted to *J. Climate*.
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- Riihimaki, L. D., M. F. Cronin, … D. Rutan,….. 2024: Ocean surface radiation measurement best practice, *Front. Mar. Sci. 11:1359149. doi: 10.3389/fmars.2024.1359149*
- Loeb, N. G., D. R. Doelling, S. Kato, W. Su, P. E. Mlynczak, and J. C. Wilkins, 2024: Continuity in top-of-atmosphere earth radiation budget observations, *J. Climate*, Accepted

Thank you for your attention!